



## Article (cont. from p. 437)

ity of the Bell Canyon, Neill et al. [1983] estimated that at most 20-50 times less salt could have been transported through the Bell Canyon than is estimated to have been removed.

The second explanation for the missing salt is provided by Bachman [1983]. According to him, major dissolution of the evaporites in the Delaware Basin has been restricted to areas where the Pecos River and its tributaries have initiated karst systems, or to limited areas which overlie the Capitan Reef aquifer. Figure 1 shows the path of an ancestral Pecos River, east of the present day Pecos, as postulated by Bachman [1983] on the basis of river gravel deposits left by this river. Bachman [1983] believes that this ancient system was responsible for the development of an extensive karst terrane now seen east of the present day river. The salt beds of the lower Salado Formation were selectively dissolved during the Cenozoic time as a result of a dissolution front which was perched on the upper anhydrite in the Castle Formation. Bachman [1983] further states that the Tertiary and Pleistocene hydrologic conditions no longer exist except along the present Pecos River channel, and therefore the probability of further dissolution in the proximity of the WIPP repository is remote.

**Mechanics of Dissolution**

The earliest proposal for a specific mechanism of dissolution of salt in the Delaware Basin was made by Lee [1925]. His mechanism of "solution and fill" postulates infiltration of rain water collected in arroyos into the fractures of soluble rock. This results in the development of a drainage system a few feet below the surface, into which surface debris is carried by subsequent storms. As the gradient of the drainage system increases, headward cutting results. According to Bachman [1980], this process is currently active in Nash Draw, west of the WIPP site.

As stated above, salt has been removed from the Rustler Formation, which lies between 170 and 260 m below the ground surface at the WIPP site. The lowest affected zone is progressively deeper to the west. In the western part of the WIPP site and in Nash Draw, the top of Salado has also been affected by dissolution, and the permeable residual thus formed contains brine at the Rustler/Salado interface. There are salt lakes in the southern part of Nash Draw and there are saline seeps along the Pecos River near Malaga Bend, about 25 km southwest of the WIPP site. These are thought to be the discharge points for the brine produced from the dissolution of Rustler halite.

A satisfactory explanation of the dissolution and removal of salt from the deeper strata is more problematical. There are at least three different schools of thought concerning the absence of halite in the lower Salado and Castle formations. On the basis of interpretation of acoustic logs from a large number of wells in the Delaware Basin, Anderson [1982] has concluded that about 50% of the salt in the Salado and Castle formations has been removed by dissolution, with as much as 70% of the original salt removed from the lower Salado horizon in the basin. For the mechanics of salt removal through "deep" dissolution, Anderson [1981] invoked a "brine density flow" model, which had been proposed earlier [Anderson and Kirkland, 1980] for the formation of breccia chimneys. This mechanism requires a connection between the lower Salado and the underlying Bell Canyon aquifer through fractures in the intervening Castle Formation. It was hypothesized that surface recharge moves into the evaporites, dissolves salt in the Salado and Castle formations, and the resulting brine sinks into the underlying aquifer. Thus, the postulated mechanism would continue as long as the supply of undersaturated water lass and the fractured pathway remains unclogged. Wood et al. [1982] studied the potential dissolution mechanisms of diffusion and convection from the halite zones of Castle and Salado to the Bell Canyon and the Capitan Reef aquifer based upon a range of reported values for the hydrogeologic and geochemical parameters which influence salt removal. They concluded that the removal of dissolved salt through the Bell Canyon can take place only at a very slow rate, which would be grossly insufficient for removing approximately  $7 \times 10^{12} \text{ m}^3$  of salt from lower Salado in 1.5 million years, or about 4.7 million  $\text{m}^3$  of salt per year (as estimated by Anderson, 1981). Even using the most conservative reported values for the hydraulic conduc-

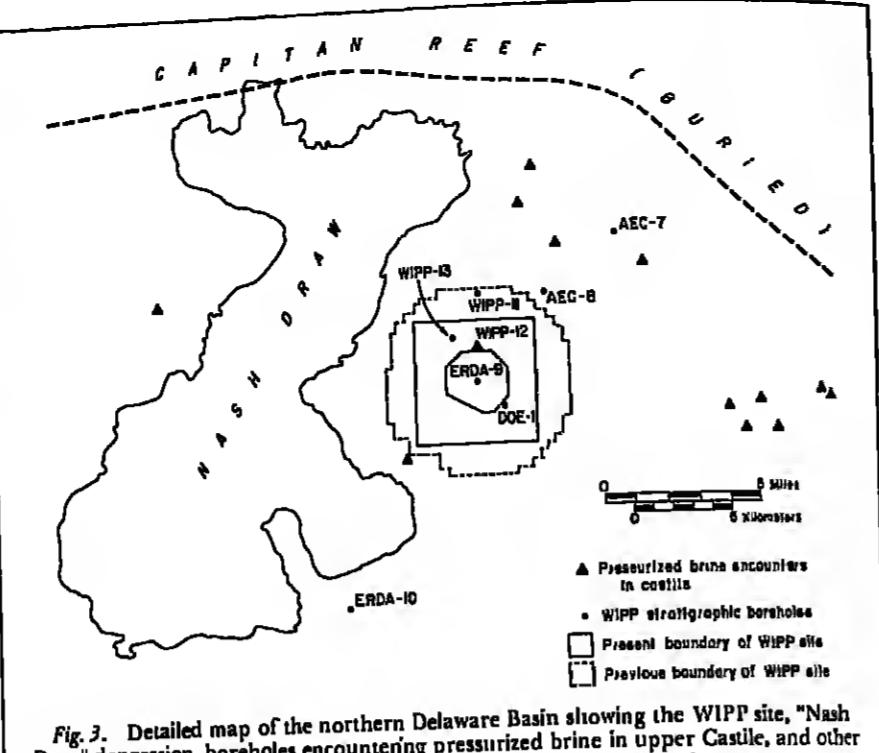


Fig. 3. Detailed map of the northern Delaware Basin showing the WIPP site, "Nash Draw" depression, boreholes encountering pressurized brine in upper Castle, and other deep stratigraphic boreholes drilled in connection with the WIPP project.

low pre-Cretaceous dissolution to reach into the Castle Formation would require a regional dip in the Cretaceous far greater than observed. Bachman [1983] believes that Tertiary and Pleistocene hydrologic conditions no longer exist along the present Pecos River channel, and therefore the dissolution of halite in the lower Salado and Castle Formations is not an active process, at least near the WIPP site.

The third explanation for the missing salt in the lower Salado is that much of the missing salt simply represents a facies change or removal during a much earlier time soon after deposition. By drawing a composite isopach map of the Castle and lower Salado Formations, Lawler [1983] has shown that the "missing halite" areas commonly result in little or no difference from regional thicknesses of Castle and lower Salado evaporites. He therefore ascribes the observation of missing salt to several factors other than Cenozoic dissolution (e.g., depositional heterogeneities, perturbation of original bed thicknesses by localized deformation, and ambiguous identification of members or marker beds).

A serious problem encountered in explaining the removal of salt through dissolution is the disposition of the resulting brine. Anderson's brine density model requires the removal of brine through the Bell Canyon aquifer; but, as mentioned above, the Bell Canyon is thought to be incapable of acting as a sink for the large amount of brine produced from dissolution. Bachman [1983] has not even addressed the question of the disposal of solution brine. Lawler [1983] also has acknowledged the difficulty in postulating a viable sink for his stratigraphic dissolution model.

### Time and Rate of Dissolution

Anderson [1981, 1982] maintains that most of the dissolution and removal of halite from the Salado and Castle evaporite beds in the Delaware Basin has occurred since the tilting of the basin to the east during the latter part of the Cenozoic, probably 4-8 my. ago. According to him, this event exposed the thick limestone of the Capitan Reef as well as the sandstones of the Bell Canyon Formation which acted as suppliers of meteoric water to the lower part of the evaporites. Anderson has also used the correlation between the area of the bulk of the missing salt with the deep depression filled with Pleistocene alluvium (described by Maley and Huffington [1983] as evidence for most of the dissolution front) to determine the age of the WIPP site from this kind of dissolution. The best way to determine the integrity of the WIPP site from this kind of dissolution is to examine the evidence front boreholes drilled at the site and surrounding it. Figure 1 shows the "deep dissolution" edges for the salt units, as interpreted by Anderson [1981]. It should be noted that the WIPP site is situated in the northern part of the basin, away from the dissolution front. The nearest point of the dissolution edge from the WIPP site is about 25 km to the southwest (Figure 1).

There are five boreholes at the WIPP site (WIPP-9, 11, 12, 13, and DOE-1) which have penetrated the lowermost anhydrite bed (Anhydrite-1) in the Castle Formation. These holes have been cored at selected intervals, and geophysical logs for the entire depth have been obtained. In addition, three holes have been obtained outside the WIPP boundary, AEC-7 and 8 to the northeast and ERDA-10 to the southwest. The boreholes were drilled, cored, and logged through the Castle Formation. None of these eight boreholes (Figure 3), and none of the several industry boreholes around the WIPP site, show any evidence of extensive dissolution. This points to the fact that at least the immediate area surrounding the WIPP site has not been affected by deep dissolution. Even if the rate of advance of "deep" dissolution was the same as the "shallow" dissolution of Bachman [1980] (i.e., 10-13 km per million years) the site appears to be safe for the next 2 million years.

A borehole located 3.25 km north of the site, drilled to assess the occurrence of potash minerals in middle Salado, encountered the Salado marker beds at elevations about 23 m below their expected occurrence. This anomalous depression has been confirmed by the logs of two other boreholes and has been suggested as a possible site of deep dissolution. The Department of Energy has accepted EEG's suggestion to drill a borehole to test the origin of this depression. This borehole, to be called DOE-2, will be drilled and cored down to the Bell Canyon Formation. The work is planned for 1984. One more anomalous feature has been pointed out near the WIPP site. About 8 km southeast of the center of the WIPP site, the acoustic log of a well shows that

bed in the lowest part of the Salado Formation may be missing. Since this well is outside the WIPP boundary and is one unconfirmed anomaly out of a large number of wells, EEG has not made any recommendations to explore this feature further.

Approximately 3000 m of underground drifts at the selected repository level 655 m below the surface have already been excavated at the WIPP site. This excavation has been conducted to validate the site characterization under the site and preliminary design validation (SPDV) program. The thickness and continuity of strata displayed in this excavation are remarkably uniform, and there is no indication of dissolution either at the repository horizon or in cores of 15 m vertical boreholes drilled from the floor and ceiling of the excavation.

### Conclusions

There are two main areas of concern with regard to the suitability of the WIPP site. The characteristics of the water-bearing zones of the Rustler Formation should be understood very thoroughly to preclude any possibility of these zones acting as pathways for migration of radionuclides to the biosphere. In addition to hydrologic testing already completed during the past 8 years, additional drilling and field testing is planned over the next 2-3 years. These forthcoming studies include the drilling and testing of several new hydrologic wells, drilling of new wells near existing hydrologic wells to convert at least eight single wells to sets of three-well for tracer tests and flow tests. The study of cores and determination of hydrologic properties from the cores as well as more geochemical testing of Rustler Formation waters are also planned. The data gained from these studies will be used to refine a hydrologic and contaminant transport model of the Rustler aquifers. In addition, a water-balance study for the site will be conducted and some suspicious depressions on the site would be bored to see

whether they are alluvial dolines. Another study will try to answer the question of the mechanism of salt removal from the Rustler Formation. These studies will help to resolve the question concerning the possibility of karst conditions in the Rustler Formation.

The second area of concern is the effect of dissolution of salt on the integrity of the repository. The Salado Formation does not appear to have been affected in and around the WIPP site by past dissolution. The suspicion of an area of point-source dissolution from below, located 3.25 km north of the center of the site, will be investigated. Although stratabound dissolution of the Rustler salt occurs across the WIPP site, such dissolution does not seem to have affected the top of the Salado Formation at the WIPP site. However, the collapsed depression of Nash Draw is only about 6.5 km west of the center of the site. It is therefore important to understand thoroughly the mechanics of removal of salt from the Rustler Formation.

The mission of WIPP calls for the permanent emplacement of transuranic waste which

would be reduced to a level of radioactivity of natural uranium ore in less than 100,000 years.

The site selected for WIPP has been characterized sufficiently to enable the analyses of worst case scenarios of breach of the repository and consequent release of radioactivity. However, a few gaps remain in our knowledge of the geologic and hydrologic characteristics of the site which relate to the transport of radioactivity to the biosphere in the event of a breach. Additional field work to close these gaps is being performed currently and will be completed before the waste is brought to the site in 1989. If the additional work indicates a possibility of release of hazardous quantities of radioactivity to the biosphere, the EEG will recommend additional engineered barriers or investigations for a new site.

**Acknowledgments**  
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The study focuses specifically on changes in hydrologic conditions such as precipitation, evaporation, soil moisture, and runoff that would accompany a twofold increase in atmospheric CO<sub>2</sub>. The Goddard model produces an atmospheric warming of 4.1°C for this doubling, which is at the high end of the range predicted by National Academy of Sciences studies (Eos, November 15, 1983, p. 929, and August 17, 1982, p. 609).

The researchers first looked at large-scale

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### News (cont. from p. 459)

ately 18 of the representatives on the House committee will meet several times during the next few months. Its first meeting is scheduled for August 2. A report will be prepared by the end of the year. House Science and Technology Committee Executive Director Harold Hanson will be coordinating the study. John Holmfield will provide staff support for the task force.

## Upcoming Hearings In Congress

The following hearing has been tentatively scheduled for the coming week by the Senate and House of Representatives. Dates and times should be verified with the committee or subcommittee holding the hearing; all offices on Capitol Hill may be reached by telephoning 202-224-3121. For guidelines on contacting a member of Congress, see AGU's Guide to Legislative Information and Contacts (Eos, April 17, 1984, p. 159).

July 31, August 1, August 2: Hearing on the status, trends, and plans for the operations management and use of the space transportation system by the Space Science and Applications Subcommittee of the House Science and Technology Committee. Rayburn Building, Room 2525, 9:30 A.M.—B7R

## Beardmore Glacier Proposals Wanted

Proposals for research projects to be conducted in the upper Beardmore Glacier area of Antarctica during the 1985-1986 field season are being accepted by the National Science Foundation (NSF) through August 16. Later proposal submissions should be discussed with the appropriate program manager (see below).

A temporary camp with helicopter support will be established in the region. Occupation by scientific parties will likely be between mid-November 1985 and mid-January 1986. Transportation in the field will be by UH 1-N twin-engine Huey helicopters (with a range of approximately 185 km) and by motor toboggans. Satellite tent camps will be established within the range of the helicopters. The exact position of the main camp will be determined in November. Likely candidates, however, are Buckley Island Quadrangle, in the area of the Walcott Névé or the Bowdoin Névé, near Coatsland Bluff or Mount Névé.

A workshop was held at the University of Maine May 25-26 on the potential for a remote field camp in the area. A report entitled, "The Beardmore Glacier Remote Field Camp," was prepared by George H. Denison of the University of Maine, and by James W. Collinson and David H. Elliot, both of The

### Geophysical Events

This is a summary of *SEAN Bulletin*, 9(6), June 30, 1984, a publication of the Smithsonian Institution's Scientific Event Alert Network. The complete bulletin is available in the microfiche edition of *SEAN* as a microfiche supplement or as a paper reprint. For the microfiche, order document E84-007 at \$2.50 (U.S.) from AGU Fulfillment, 2000 Florida Avenue, N.W., Washington, DC 20009. For the paper reprint, order *SEAN Bulletin* (giving volume and issue numbers and issue date) through AGU Separates at the above address; the price is \$3.00 for one copy of each issue number; for those who do not have a deposit account, \$2 for those who do; additional copies of each issue number are \$1. Subscriptions to *SEAN Bulletin* are available from AGU Fulfillment at the above address; the price is \$18 for 12 monthly issues mailed to a U.S. address, \$26 if mailed elsewhere, and must be prepaid.

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June, 85 sites (50%) were in the normal range, and only 19 sites (11%) were well below average.

Although most of the northern United States experienced a wet June, low flow persisted in Texas, and dry conditions developed in much of the southeast and in the Ohio River Valley.

Reflecting overall conditions in June, the combined average flow of the nation's three major rivers was 1,184 billion gallons per day (38% above the seasonal average). The Mississippi, St. Lawrence, and Columbia rivers together drain more than half of the lower 48 states, and their flows provide a convenient check on the status of the nation's water resources.

Record high monthly average flows for June occurred in Iowa, Kansas, Minnesota, Nebraska, and Nevada. The Humboldt River at Palisade, Nev., for example, set a record for June of almost 9 bfg, the highest flow in 77 years of record. The Humboldt River has been in the above-normal range—within the highest 25% of historic record—now for 24 straight months.

Record low or near record low flows were recorded in Hawaii, Idaho, Louisiana, and Texas. Extremely dry conditions prevailed in parts of Texas, with zero average flow of the North Concho River near Carlsbad, the lowest in 37 years of record. Barely 200,000 gallons per day of flow were recorded on the North Bosque River near Clifton, Tex., the lowest June flow in 61 years of record.

Flows of five of the nation's largest rivers for June—Mississippi River at Vicksburg, Miss., 598 bfg, 69% above average, despite a 38% decline from the average May flow; the Columbia River at The Dalles, Ore., 386 bfg, 17% above the long-term average for June and up 84% over last month; the St. Lawrence River near Massena, N.Y., 200 bfg, 10% above average, about the same flow as May; the Missouri River near Hermann, Mo., 137 bfg, 146% above average and up 5% from May; and the Ohio River at Louisville, Ky., 53 bfg, 31% above the June average, although flow declined by 63% from May. (Map courtesy of USGS, Reston, Va.)

The astronomers arrived at this explanation for the star's behavior after they detected periodic changes in the shape and symmetry of its H-alpha spectral signal, changes conforming to the short rotation period. The changes were greater than what could be explained solely on the basis of the velocity of the star's surface, however. The astronomers concluded from the alternating redward and blueward shifts and from other evidence that the emitted material forms a stellar wind with oppositely directed streams whose high velocities cause the excessive wavelength shifts. —PR

## June Streamflow

Flows of the nation's key index gauging streams were at average to well-above average levels for June across much of the country, according to a regular check on the condition of the nation's water resources by the U.S. Geological Survey (USGS).

USGS hydrologists said that of the 170 key index gauging stations nationwide, flows at 66 sites (39%) were well above average for

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